Implications of accounting for land use in simulations of ecosystem services and carbon cycling in Africa

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Abstract

Dynamic global vegetation models (DGVMs) are important tools for modelling impacts of global change on ecosystem services. However, current models often do not take full account of human land management and land use and land cover changes (LULCC). We integrated croplands and their management and natural vegetation recovery and succession following land use abandonment into the LPJ-GUESS DGVM. The revised model was applied to Africa as a case study to investigate the implications of accounting for land use on agricultural production, net ecosystem carbon balance (NECB) and on the general skill of the model in reproducing trends and patterns in vegetation structure and function. The seasonality of modelled monthly fraction of absorbed photosynthetically active radiation (FPAR) was shown to agree well with satellite-inferred normalised difference vegetation index (NDVI). In regions with a large proportion of cropland, the managed land addition improved the FPAR vs. NDVI fit significantly. Modelled 1991–1995 average yields for the seven most important African crops, representing potential optimal yields limited only by climate forcings, were generally higher than reported FAO yields by a factor of 2–6, similar to previous yield gap estimates. Modelled inter-annual yield variations during 1971–2005 generally agreed well with FAO statistics, especially in regions with pronounced climate seasonality. Modelled land-atmosphere carbon fluxes for Africa associated with land use change (0.09 PgC yr⁻¹ release to the atmosphere for the 1980s) agreed well with previous estimates. Crop-land management options (residue removal, grass as cover crop) were shown to be of similar importance to the land-atmosphere carbon flux as land use change for the 20th century.

1 Introduction

A large part of the earth’s surface is covered by managed land, e.g. croplands (about 11%) and pasture (about 24%) (Klein Goldewijk et al., 2011). Increased demand for
food, fiber and energy production by a growing human population is responsible for an expansion of managed land into intact natural vegetation and more intensive management of existing managed land. This is often in conflict with the goal of preserving global biodiversity (Heywood, 1995). Anthropogenic land use and land cover changes (LULCC) have been important contributors to rising atmospheric CO$_2$ concentrations [hereinafter “CO$_2$”] in the past, accounting for an estimated 156 PgC released for the period 1850–2000 (Houghton, 2003). Land use, particularly tropical deforestation, continues to contribute significantly to anthropogenic greenhouse gas emissions, releasing an average of around 2 PgC yr$^{-1}$ for the 1980s and 1990s (Houghton, 2003; Le Quéré et al., 2009; Zaehle et al., 2011). LULCC may also lead to changes in land surface properties and near-surface energy, water and momentum exchange, contributing to biophysical climate forcing (Pitman et al., 2009; Wramneby et al., 2010; Arora and Montenegro, 2011).

Climate change is predicted to cause adverse effects on terrestrial vegetation in large parts of the world, putting further stress on both biodiversity and managed land productivity (Foley et al., 2005). Combined effects of rising CO$_2$, climate change and LULCC will act to influence the role of the terrestrial biosphere in the global carbon budget by altering net primary productivity (NPP) and soil carbon turnover, altering the potential of soil and vegetation to store carbon (McGuire et al., 2001).

Agricultural production of food and fodder is one of the key ecosystem services underpinning the well-being and development of human society (Millennium Ecosystem Assessment, 2005). Increasing temperatures are projected to have negative consequences for crop yields in many regions, especially in the lower latitudes that already experience warm temperatures. Changes in precipitation will be more variable regionally and might exacerbate or alleviate the negative effects of raised temperatures, depending on the sign of change (Roudier et al., 2011). Furthermore, an increase in climate variability leading to stronger and more frequent extreme events, such as droughts and floods, may generally be expected to cause a reduction of crop yields (Rowhani et al., 2011). A CO$_2$ fertilisation effect may reduce the negative effects of
raised temperatures on crop production but the overall magnitude and persistence of such an effect is still debated (Ainsworth and Long 2005; Long et al., 2006; Tubiello et al., 2007; Müller et al., 2009).

Africa has been identified as one of the world’s most vulnerable regions to climate change (Müller et al., 2011). Predicted decreasing crop yields combined with a rapid population increase make food shortage a likely scenario (Müller et al., 2009), but the pattern is regionally complex. A review of both empirical and processed-based crop models on future yields in West Africa shows a high degree of agreement on the negative crop yield prospects (Roudier et al., 2011). In East Africa, precipitation is predicted to increase in the future according to projections from a majority of general circulation models (GCMs). A resulting possible increase in NPP may also lead to increased crop yields (Doherty et al., 2010).

The African continent has been acting as a carbon sink for most of the 20th century, sequestering carbon mainly into forest biomass (Ciais et al., 2009). Intact African tropical forests stored 0.34 PgC yr$^{-1}$ during the past 40 yr (Lewis et al., 2009) while in the absence of fire, CO$_2$ uptake in African savannahs has been estimated at 1.9 PgC yr$^{-1}$ (Bombelli et al., 2009). Savannah fires, either as natural episodic events or resulting from human management, release much of this carbon back to the atmosphere (van der Werf et al., 2008; Lehsten et al., 2009). LULCC, such as the conversion of forest into cropland or pasture, releases large amounts of CO$_2$ and also reduces the carbon sink potential. The present-day African deforestation-related carbon loss has been estimated at 0.4 PgC yr$^{-1}$ (Williams et al., 2007), which is relatively modest compared with other tropical regions, mainly because of lower deforestation rates compared to other tropical forest regions. Conversely, abandonment of cropland and the subsequent succession into forest sequesters atmospheric CO$_2$, especially during early successional phases (Houghton et al., 1983). Improved knowledge of carbon fluxes related to deforestation, afforestation, and crop and pasture processes may provide important information to help determine the preferred land management options for an area with respect to carbon sequestration.
Dynamic global vegetation models (DGVMs) have been used for modelling terrestrial vegetation dynamics and ecosystem biogeochemistry under changing climate and CO₂, and have been widely used to investigate the potential for terrestrial biota to feed back to climate change, particularly through carbon cycle changes (McGuire et al., 2001; Friend et al., 2007; Sitch et al., 2008; Arneth et al., 2010a, b). At first DGVMs accounted for LULCC by simulating grassland instead of natural forest or by harvesting a fraction of the natural NPP (McGuire et al., 2001; Brovkin et al., 2004). However, the large variety of crop management practices worldwide made it clear that the implementation of specific crop or pasture functions (phenology, carbon allocation and overall productivity) within DGVMs was required. Several DGVMs now have the capabilities of accounting for both natural vegetation and agriculture (crops and/or pasture), either at the scale of a large region (Kucharik and Brye, 2003; de Noblet-Ducoudré et al., 2004; Gervois et al., 2004; Kucharik and Twine, 2007; Van den Hoof et al., 2011; Berg et al., 2011; Hidy et al., 2012) or globally (Bondeau et al. 2007, Müller et al., 2007).

To evaluate the importance of croplands and pasture for land-atmosphere coupling and the carbon cycle, we integrated managed land and land use change functionality into a DGVM optimised for regional applications, further developing an approach adopted from Bondeau et al. (2007). Focusing on Africa as a case study, we simulated crop production, seasonal and interannual variability of vegetation productivity and the continental-scale land-atmosphere carbon fluxes forced by recent historical climate, CO₂ and observed land use and land cover. We evaluate the skill of the model in comparison to observations and investigate how the introduction of managed land affects the predictions of the model across the climate space of Africa.
2 Methods

2.1 DGVM

We employed LPJ-GUESS (Lund-Potsdam-Jena General Ecosystem Simulator, Smith et al., 2001), a DGVM optimised for regional applications, as the modelling platform for our study. LPJ-GUESS combines a detailed, individual- and patch-based representation of vegetation structure, demography and resource competition with physiological and biogeochemical process formulations in common with the more generalised global model LPJ-DGVM as detailed by Sitch et al. (2003) with modifications by Gerten et al. (2004). The model was applied on a 0.5 × 0.5° regular latitude-longitude grid across Africa. The potential natural vegetation in each grid cell was simulated as a dynamic mixture of up to six tree and two grass plant functional types (PFTs), using the PFT set and parameters described in Ahlström et al. (2012a). Photosynthesis, respiration, tissue turnover and carbon allocation to leaves, fine roots and stems (sapwood) are modelled independently for each tree, or generically for C₃ and C₄ grasses, respectively, in each of a number (20 in this study) of replicate patches accommodating variation of vegetation dynamics within the potential natural vegetation fraction of each simulated grid cell. Height and diameter growth of trees are regulated by carbon allocation, sapwood-to-heartwood conversion, and a set of prescribed allometric relationships for each PFT. Population dynamics (establishment and mortality) are influenced by current availability and uptake of light and soil water, stand size structure, and the life history characteristics of each PFT (Hickler et al., 2004). Biomass-destroying disturbances are simulated as a stochastic process, here with a generic expectation of 0.01 yr⁻¹. In addition, fires are modelled prognostically based on temperature, current fuel load and moisture (Thonicke et al., 2001). A detailed description of LPJ-GUESS is given by Smith et al. (2001).

LPJ-GUESS has been evaluated extensively and exhibits comparable skill to other approaches and models in reproducing observed temporal and spatial variation in large-scale vegetation patterns, productivity and ecosystem-atmosphere carbon
exchange (Piao et al., 2013). Results from the model have been compared to ecosystem flux measurements, site measurements and satellite-based proxies of NPP, leaf area index (LAI) and biomass, spanning many of the world’s biomes (Morales et al., 2005; Hickler et al., 2005, 2006; Hély et al., 2006; Smith et al. 2008; Ahlström et al., 2012a).

Land use and cropland representations were implemented in LPJ-GUESS based on LPJ-mL (Bondeau et al., 2007), with a number of modifications (see below). Cropland is represented by eleven crop PFTs (Table 1), simulated separately (without inter-PFT competition) and two grass PFTs (C$_3$ and C$_4$ grass) as cover crop between harvest and sowing. The same grass PFTs are used to represent pastures. Irrigated crops are simulated separately. For wheat, an upper temperature limit of 10°C for the coldest month for growth is set following Bondeau et al. (2007). The main modifications in the model compared to LPJ-mL as described by Bondeau et al. (2007) are a new phenology scheme, coupling LAI and leaf C mass on a daily timestep, a dynamic PHU calculation based on local climate conditions in each grid cell, simulation of regeneration and recovery of vegetation by creating a new natural vegetation stand following cropland abandonment and a revised calculation of crop sowing dates following Waha et al. (2011).

In LPJ-mL, crop phenology followed the development of the potential LAI, which was coupled to the sum of heat units (degree days above a crop-specific base temperature, $T_b$) accumulated from the time of sowing, and eventually corrected if a climatic stress reduced the actual available carbon for leaf growth (Bondeau et al., 2007). This could cause significant discrepancies between daily values of heat sum-based LAI and actual available carbon for leaf growth based on the simulated leaf carbon assimilation. This method also requires a fixed maximum LAI value that does not respond to a changing environment, e.g. changing CO$_2$. In LPJ-GUESS, we introduced instead a feedback between daily leaf carbon mass and LAI, making the leaf development, before flowering and the onset of senescence, fully consistent with physiological constraints (Fig. 1). Upon sowing, the initial carbon was set to 10 g m$^{-2}$.
The development of the harvest index (HI), i.e. the fraction of above-ground carbon present in the harvestable organs, and the onset of senescence and LAI development during senescence is dependent on accumulated heat units and calculated as in Bondeau et al. (2007). In LPJ-GUESS, the potential heat unit (PHU) sum needed for full development of a crop, determining the time at which the crop is harvested, is calculated dynamically, using a 10-yr running mean of heat unit sums accumulated from the sowing date to the end of a sampling period (ranging from 190 to 245 days) derived from default sowing and harvest limit dates as reported by Bondeau et al. (2007). The dynamic PHU calculation can be done either for an initial time period only, to calibrate for the local climate, or also for an extended period, to simulate adaptation to a changing climate by selecting suitable crop varieties/genotypes. In the present report, dynamic PHU was used for the whole time period (Table 1). A lower PHU limit of 900 degree days was used. Maintenance respiration of storage organs was set to zero. The modelled crop yield represents potential optimal yield, limited by climate and CO₂ only.

The recent local climate variability (seasonality) determines whether the limiting factor for date of sowing is temperature or precipitation, or if there is a lack of climate seasonality (Waha et al., 2011) (Fig. 2). Five main seasonality types to determine sowing date were specified as: TEMP (temperature seasonality); PREC (precipitation seasonality); TEMPPREC (both temperature and precipitation seasonality, minimum monthly temperature < 10°C), temperature determines sowing date; PRECTEMP (both temperature and precipitation seasonality, minimum monthly temperature = 10°C), precipitation determines sowing date; NONE (no seasonality), default sowing date used. For irrigated crops at PRECTEMP sites, temperature-dependent sowing was used unless specifically stated in the text. The temperature limits for temperature-dependent sowing were as in Waha et al. (2011) (Table 1).

Croplands were harvested each year (see Table 2 for harvest parameters used in this study). A PFT-specific fraction of the harvestable organs (the harvest efficiency) constitutes the yield (multiplied by 2.0 for deriving the total dry yield from carbon units) and is assumed to be oxidised within a year. Of the above-ground carbon left after harvest, a
further fraction (the residue removal fraction) is removed (and oxidised within one year). This PFT parameter is set to 0.75 for all crops in this study unless specified otherwise. Grazing of managed grassland is simulated by removing 50% of the above-ground carbon. This is roughly in agreement with estimates of a 90% removal in intensively grazed pastures and a 50% re-entering of this carbon back to the litter pool as manure.

At the conversion of forest to cropland, 70% of tree stems are harvested and the rest is oxidised the same year (burned). Part of the harvested wood (firewood) is oxidised the same year (67%) and the rest goes to a slow pool, representing paper and timber, with a 25-yr turnover period. This is a simplified version of Houghton’s et al. (1983), where two long-lived pools (10 yr/100 yr turnover times) exist, also used by McGuire et al. (2001) with LPJ. At cropland abandonment (conversion to natural vegetation), a new stand is created from bare ground, allowing an establishment of natural vegetation with a succession e.g. from grass to deciduous trees and finally evergreen trees. An example of modelled carbon fluxes at cropland expansion and abandonment is shown for a moist tropical site in Liberia in Fig. 3.

“Land use functionality” in this paper refers to the dynamic representations of different landcover types with associated land uses, in this study specifically irrigated or rainfed cropland and pasture. The expected difference in e.g. modelled carbon fluxes in simulations with and without land use functionality would be the differential response by different land covers to changes in climate and CO$_2$ as well as the effects of LULCC.

2.2 Observational data

Soil texture data were as described by Sitch et al. (2003) and historical yearly global atmospheric CO$_2$ concentrations were obtained from the Carbon Cycle Model Linkage project (McGuire et al., 2001). Monthly values of temperature, precipitation and cloudiness were taken from the CRU TS 3.0 data set at 0.5 degree resolution (Mitchell and Jones, 2005), provided by the British Atmospheric Data Centre (BADC), National Centre for Atmospheric Science, Natural Environment Research Council. Available from http://badc.nerc.ac.uk/view/badc.nerc.ac.uk_ATOM_dataent_1256223773328276.
We used the bimonthly maximum value composited Normalised Difference Vegetation Index (NDVI) Global Inventory Modeling and Mapping Studies (GIMMS) dataset derived from the National Oceanic and Atmospheric Administration’s (NOAA) series of Advanced Very High Resolution Radiometer (AVHRR) instruments with a spatial resolution of 0.07° (Tucker et al., 2005). The red and near infrared bands of the AVHRR sensor are used to compute the NDVI, a metric that is correlated with the photosynthetic activity and LAI of green vegetation (Sellers, 1989). The NDVI also reduces variations not related to vegetation information from the effects of topography, sensor calibration and viewing geometry, while the maximum value compositing procedure minimises the effects of cloud cover and the atmosphere (Holben, 1986). Corrections for sensor differences, sensor drift and the effects of volcanic eruptions have all been accounted for in the GIMMS data set. We scaled up the NDVI observations to 0.5° using mean-value aggregation, by including all overlapping cells within the larger 0.5° area and by ignoring all water areas, negative values, and urban land cover (identified using the Global Land Cover Facility land cover dataset, e.g. Hansen et al., 2000). We also extracted the maximum NDVI from the two NDVI estimates for each month over the whole time series in order to match the temporal frequency of LPJ-GUESS output. This has the added benefit of further reducing the effects of atmospheric disturbance in the data.

Food and Agriculture Organization of the United Nations (FAO) national crop yield data were obtained from http://faostat.fao.org/site/567/default.aspx{#}ancor(last access: 1 March 2012).

The historical cropland data set used in this study was an adaptation by Bondeau et al. (2007) of the cropland fraction for the period 1901–1992 (Ramankutty and Foley, 1999), the distribution of crops for 1992 (Leff et al., 2004), the pasture fraction for 1970 (Klein Goldewijk and Batjes, 1997) and the irrigated agricultural fraction for 1995 (Döll and Siebert, 1999).
2.3 Model experiments and evaluation

Model spin-up was for 300-yr, applying the year 1901 CO$_2$ value and temperature-detrended climate data for 1901–1930 in repeated cycles. In the experiments with land use functionality turned on, the 1901 cropland and pasture fractions were used for the spinup period. Cropland and pasture stands contained only one patch since stochastic processes of natural vegetation establishment and mortality, or disturbances by fire or other events, are not needed for managed ecosystems.

To test the accuracy of the model to capture the seasonal vegetation greenness variation of the African continent, with and without land use functionality, modelled monthly mean Fraction of absorbed Photosynthetically Active Radiation (FPAR) was compared with observed monthly maximum NDVI for the period 1982–2006. Although FPAR and NDVI are not comparable in terms of their absolute values, they can both serve as indicators of seasonally varying phenology and leaf growth dynamics. Twelve sites located in the major African crop-intensive regions, representative of the different climate seasonalities, crop types and irrigation intensities, were selected to provide a significant crop-derived component to the total FPAR/NDVI levels (Table 3, Fig. 2).

Standardised anomalies (zscores) of FPAR and NDVI were calculated as deviations from the population mean, divided by the standard deviation. The population was all the values of monthly data either for a given site or for the entire African continent, the latter for comparisons of seasonal fluctuation amplitudes. Distances between standardised FPAR and NDVI anomalies were computed by the distance index for a site, where $N = \text{number of monthly FPAR/NDVI data pairs}$ and $\Sigma$ is the sum over all these data pairs (Eq. 1).

$$d_i = \sqrt{\frac{\sum (\text{zscore(FPAR)} - \text{zscore(NDVI)})^2}{N}}$$

A number of experiments were performed over the period 1901–2006 to investigate the sensitivity of net ecosystem carbon balance (NECB) of the African continent to LULCC
These experiments were designed to investigate separately the relative contributions of climate change, changes in CO$_2$ and management practices. Factorial simulations were performed, keeping CO$_2$, land use, net climate or individual climate drivers (temperature, precipitation, incoming shortwave radiation) at the spinup levels (see above) while varying other drivers, and compared to a simulation with all drivers varying.

3 Results and discussion

3.1 Seasonal variation of vegetation greenness

Averaged over 10–15° latitude bands covering the African continent, seasonal anomalies of modelled FPAR generally matched remotely sensed NDVI well, although sometimes preceding the NDVI variation by 1–2 months (Fig. 4). In large parts of Africa, the modelled phenology is governed directly by precipitation (Fig. 2), so the accuracy of the climate data (monthly in the present study) is critical and could be a source of error. In Africa, the density of the climate sampling network is relatively sparse in many places, and this can introduce a bias in gridded climate products, especially in precipitation. A bi-seasonal pattern in NDVI evident around the equator is accurately captured by the model. An exception to the good FPAR vs. NDVI fit is seen in the 15–25° N latitude band largely characterised by desert with only little vegetation. The data-model mismatch could reflect either NDVI artefacts caused, e.g. by seasonal atmospheric dust clouds, or poor seasonal accuracy in the climate data for this region. Adding land use functionality to LPJ-GUESS brings rather minor changes to these continental-scale seasonal FPAR patterns, reflecting the relatively small cropland fraction of Africa, which when averaged over the latitudinal bands does not exceed 9% of the area (Ramankutty and Foley, 1999). Moreover, since most crop areas and pastures in Africa are not irrigated, the greening of natural vegetation and managed land
follows rather similar patterns over the year, since in the largest part of the continent precipitation is a major determinant of crop production (Fig. 2).

When focusing on a number of locations in regions with a relatively large fraction of cropland (Table 3, Fig. 2), the difference in leaf area development over the course of a year between managed land and natural vegetation is evident (Fig. 5). The most apparent deviations of LPJ-GUESS potential natural vegetation seasonal FPAR from observed NDVI are in regions where crops have replaced a natural vegetation of evergreen or mixed forest, having little or no seasonal variation. Including land use functionality improves the fit in these areas. In our examples, this is the case in temperate northern (Algeria) and southern Africa (South Africa 1), subtropical East Africa (Ethiopia) and tropical East Africa (Uganda and Tanzania).

At savannah locations (examples are for Senegal, Sudan and Burkina Faso) the growing season onset simulated as an increase in modelled natural vegetation FPAR is 2–3 months ahead of the increase in NDVI. Accounting for land management in the model partially reduces this difference, since the crop FPAR increase is significantly delayed compared to natural vegetation and, at these locations, quite similar to the NDVI increase. In regions where natural vegetation is dominated by grasslands, including land use has little influence on the FPAR vs. NDVI fit. This is with the exception for the region in Egypt, where the modelled rainfed natural vegetation FPAR is sparse and deviates highly from the observed NDVI levels and seasonal pattern. Two distinct growing periods are apparent in the NDVI seasonality, reflecting the irrigated double-cropping farming practice in the Nile delta. When including irrigated cropland in the model (Fig. 5), the maximum FPAR is increased to realistic levels but at present, the model can reproduce only a single growing season. Double-cropping capability is a high priority in future versions of the model.

Altogether, simulations including land use improves the FPAR vs. NDVI fit in the most crop-intensive regions across the entire continent, with the exception of the Nile valley and delta (Fig. 6).
3.2 Inter-annual greenness variability 1982–2006

The Sahel region has been well studied with respect to recent climatic and land cover trends. Hickler et al. (2005) showed that a positive trend in satellite-measured NDVI greenness through the 1980’s and 1990’s was best explained by a positive trend in precipitation over this period. Expanding this analysis to 2002 and using a spatially-explicit approach, accounting for additional regional effects of human land use, did not significantly improve agreement between modelled vegetation patterns and satellite observations (Seaquist et al., 2009). Anomalies of NDVI and FPAR simulated by our model, averaged over the Sahel rectangle (12–20° N, 19° W–40° E) for the period 1982–2006, are shown in Fig. 7. Similarly to the previous studies cited above, an overall positive trend with correlated interannual variation in FPAR and NDVI are apparent and may be explained by variation in rainfall. After the late 1990s, the trend levels off, as observed before (Kaspersen et al., 2011). Inclusion of land use had only a minor effect on simulated FPAR, reflecting the low (ca. 5 %) cropland cover in the region.

3.3 Comparing modelled crop yields with FAO statistics

In its present configuration LPJ-GUESS aims to reproduce crop seasonal phenology as well as potential, rather than average observed yield. Modelled crop yield thus represents an upper limit for a particular location given by climate and CO₂, not reduced by such factors as nutrient limitations, crop pathogens, herbivory, weeds or suboptimal farming practices. The difference between potential optimal yield and actual yield, often termed the yield gap, varies between different regions and countries. Large parts of Africa have the largest yield gaps in the world (Lobell et al., 2009; Licker et al., 2010). A comparison of modelled yields for the seven most important crops in Africa (by area), with FAO country-level crop yield statistics for the 1991–1995 period (Fig. 8) generally reflects this with higher modelled than observed yields, but for some crops, this relationship is less pronounced. There is also a variable degree of correlation between modelled and reported yield (Fig. 8).
Modelled yields of the two major African crops, maize and millet, were almost always higher than reported yields, typically by factors ranging from 2-6. In previous yield gap assessments for maize in Africa, this factor ranged from 2 to 8 (Pingali and Pandey, 2001; Lobell et al., 2009). Estimates from western Kenya of potential and actual maize yields were 3.7 and 1.7 ton ha\(^{-1}\), respectively (Tittonell et al., 2008) while our results show modelled vs. reported maize yields for Kenya of 3.4 and 1.8 ton ha\(^{-1}\), respectively. Combining statistical, GIS, socio-economic and methodology of agroecological zones for large regions of the continent, You et al. (2009) have estimated 3.5–5 ton ha\(^{-1}\) as potential yields from rainfed, high-input maize. Modelling potential optimal millet yields in the Sahel region with the ORCHIDEE DGVM, Berg et al. (2011) obtained 2–5 times higher yields than FAO reported yields. For pulses, the relationship between modelled vs. reported yields could be separated into different regions, with one group consisting of Sahel countries with a low modelled yield of up to 1 ton ha\(^{-1}\), a second group containing the majority of countries (mostly tropical) with a modelled yield of 2–4 ton ha\(^{-1}\), and a third small group of countries (Morocco, Tunisia, Egypt and Lesotho), with a modelled yield of 7–9 ton ha\(^{-1}\) (Fig. 8). With the exception of Egypt, probably because of irrigation, the reported yield was only up to 1 ton ha\(^{-1}\), resulting in large differences in estimated yield gaps for the three groups.

Reported yields exceeded modelled yields for some crops (e.g. peanut) in a few regions, e.g. the Sahel, possibly reflecting poor yield statistics or climate data for that region (see the discussion on FAO statistics in the Sahel in Berg et al. (2011)). Modelled low or zero yield for wheat is explained by the set upper temperature limit for wheat cultivation in the model (see Methods). Cassava shows a very poor correlation between modelled and reported yield. Cassava is mainly grown in regions with little climate variability, suggesting that non-climatic factors, not represented by our model, may explain the reported variation in reported yields.

Interannual variability of simulated and reported yields is a further indicator of model performance. Simulated variation in maize yields for the period 1971–2005 shows acceptable general agreement with observed yields (Fig. 9).
In addition to the known uncertainties in climate input and reported yield statistics (Hall, 1984), irrigation fractions are currently only rough estimates for each crop (cf. Bondeau et al., 2007). Unsurprisingly, irrigation can increase simulated yields substantially, but modelled rainfed crop yields also exceed reported yields in most countries (Fig. 9a). The irrigated area in Africa is well below 5% of the crop area (You et al., 2009), but whether large scale irrigation will be feasible in African regions as an adaptive measure to mitigate climate change impacts depends on future economic development (Liu et al., 2008; Ziervogel et al., 2008).

### 3.4 Modelling carbon fluxes of natural vegetation, pasture and croplands

NECB simulated for 1901–2006 with and without representations of cropland, pasture and land use change and under alternative management options – expected to alter the carbon balance – is shown in Fig. 10. Potential natural vegetation was estimated to accrue 29.2 PgC from the atmosphere during this period for Africa as a whole. When land use was accounted for, this sink was reduced to 20.9 PgC. Land use change accounted for a 4.4 PgC flux to the atmosphere (Fig. 11a), while the rest of the reduction (3.9 PgC) was due to the differential carbon uptake of the simulated potential natural vegetation and managed land (at the 1901 cover). Alternative cropland management options had a large effect on the carbon balance, as has previously been suggested e.g. by Lal (2002). When 100% crop residue removal was assumed, the carbon sink was reduced by 1.8 PgC compared to no removal. Grass cover crop growth added 1.6 PgC to the sink compared to no cover crop (Fig. 10). In our model, alternative cropping practices thus had the potential to influence the biosphere-atmosphere carbon balance to a comparable degree to land use change, underlining the need for valid characterisation of land management practices in models as well as good estimates of land cover change per se.

Model-based studies of transient changes in the global terrestrial ecosystem carbon balance over the 20th century indicate a switch from a neutral or source to a sink state around 1960 as a result of saturation of land use expansion, accompanied by
increased greenhouse gas emissions, raising the production enhancement of vegetation due to atmospheric [CO₂] (McGuire et al., 2001; Le Quéré et al., 2009; Ahlström et al., 2012b). Similar global transient responses were obtained by Bondeau et al. (2007). In our simulation, the African continent turned from being neutral to being a carbon sink already around 1950 (Fig. 10). The earlier transition to carbon uptake for the African continent, compared to global patterns, may be attributed to the continent’s lower deforestation rates in the 20th century, compared to other regions (e.g. South America) (Ciais et al., 2009). Short episodes of carbon release in the 1940s and 1980s correlated with periods of relatively warm and dry climate, conditions under which reduced plant production tends to coincide with increased heterotrophic respiration and wildfire biomass burning in the model (Fig. 11b).

The modelled African continent NECB averages for the periods 1980–1989, 1990–1999 and 2000–2005 were 0.17, −0.37 and −0.80 PgC yr⁻¹, respectively, when land use was accounted for. The change from a carbon source in the 1980s to a sink in the 1990s for African vegetation was also seen by Ciais et al. (2009), modelling land use change using the vegetation model ORCHIDEE (0.14 and −0.15 PgC yr⁻¹, respectively). The land use change component of the NECB for the 1980s was 0.09 PgC yr⁻¹. Previous studies using the same historical cropland fraction data as in this study (Ramankutty and Foley, 1999), with two different vegetation/carbon cycle models, produced similar land use change-derived carbon fluxes for the 1980s, 0.08 PgC yr⁻¹ for tropical Africa (Jain and Yang, 2005) and 0.08 PgC yr⁻¹ for the continent as a whole (Ciais et al., 2009). A book-keeping model of vegetation-atmosphere carbon fluxes with land use change input based on deforestation rates produced either similar (0.09 PgC yr⁻¹ for 1984–1990, DeFries et al., 2002) or higher (0.3 PgC yr⁻¹ for 1980–1989, Houghton, 2003) results for tropical Africa, depending on the input data. The land use input data, rather than the model used, have been implicated as the main factor explaining the differences in carbon balance estimates obtained by different studies (McGuire et al., 2001; Jain and Yang, 2005). Bondeau et al. (2007) showed that using the IMAGE land use data (IMAGE team, 2001) instead of the Ramankutty
and Foley (1999) land cover data to drive LPJ-mL, resulted in global carbon fluxes much more in agreement with estimates based on book-keeping (Bondeau et al., 2007; Houghton 2003).

With respect to geographic location, African regions with a large fraction of managed land generally had similar 1961–1990 average annual NECB compared to potential natural vegetation, but some cropland-intensive areas in East Africa had NECB-values up to 0.28 kg C m$^{-2}$ yr$^{-1}$ higher with land use taken into account (Fig. 12a). Part of this is due to a land use change-derived CO$_2$ release to the atmosphere (up to 0.22 kg C m$^{-2}$ yr$^{-1}$) (Fig. 12b). CO$_2$ capture by abandoned cropland in Africa during this period is minimal, since this area is very small.

The results from factorial driver simulations, shown in Fig. 11, suggest that CO$_2$ fertilisation had a greater influence on African NECB ($-32.5$ PgC for the period 1901–2006 than climate change ($+8.5$ PgC), land use change ($+4.4$ PgC) or alternative cropland management (up to $2.9$ PgC). However, there are many discussions on the CO$_2$ fertilisation effect (Long et al., 2006; Tubiello et al., 2007), leading to some uncertainties on the capabilities of DGVMs to correctly account for it. For the 1980s, the relationship between CO$_2$ and land use change was similar ($-0.5$ and $+0.09$ PgC yr$^{-1}$ respectively), but the relative contribution of climate was much higher ($+0.6$ PgC yr$^{-1}$). Separating the climate components, temperature and precipitation both contributed to a net flux to the atmosphere when synchronised, as in the warm and dry 1980–1996 period. After 1996, precipitation shows a positive trend, resulting in a net carbon uptake by the vegetation, overriding the negative, evapotranspiration-mediated effect of rising temperatures on ecosystem productivity.

4 Conclusions

We demonstrate the implications of introducing dynamic land management and land cover changes to a process-based dynamic ecosystem model in simulations over Africa. We showed that the correspondence between modelled and observed seasonal...
variation of vegetation cover in African regions with a high cropland area fraction was improved by accounting for land use and that the modelled crop yields, representing potential optimal yields, were in most cases in line with reported yields and corresponding yield gaps, when these were available. Land-atmosphere carbon balance estimates over Africa were changed significantly by introducing land use to the model, suggesting that the anthropogenic influence on vegetation cover and dynamics is important to consider in studies of the regional carbon balance, even in this region where anthropogenic impacts on landscapes are arguably only moderate by world standards.

The results demonstrate the applicability of land use-enabled DGVMs to assessing the individual and combined impacts of climate, management and land use policy on regional carbon balances and ecosystem services such as agricultural production. As a process-based model with a uniquely detailed implementation of vegetation dynamics, LPJ-GUESS should prove useful particularly for applications under scenarios of the future world, in which environmental conditions without modern analogues may break empirical relationships between drivers and outcomes of ecosystem functioning.

Improvements to the model currently under development include soil and vegetation nitrogen cycling and N-limits on plant production, which should benefit the modelling of both crop production and vegetation succession of abandoned cropland, especially on tropical soils with low nutrient status. For applications to Africa, refined land cover and climate (especially precipitation) input data may be expected to improve the accuracy and utility of the model.

Acknowledgements. This study was funded by the Foundation for Strategic Environmental Research (Mistra) under the Mistra-SWECIA programme and by the Formas Strong Research Environment Land Use Today and Tomorrow. The study is a contribution to the European Union Seventh Framework Programme project ClimAfrica (FP7-ENV-2009-1, 244240) and to the strategic research areas Modelling the Regional and Global Earth System (MERGE), Biodiversity and Ecosystem Services in a Changing Climate (BECC) and the Lund Centre for Carbon Cycle and Climate Interactions (LUCCI). We thank J. Ardö for help in processing the NDVI data.
References


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Table 1. PFT growing season parameters for crop PFTs used in this study.*

<table>
<thead>
<tr>
<th>Crop PFT</th>
<th>$T_{sow}$</th>
<th>$T_b$</th>
<th>PHU range</th>
</tr>
</thead>
<tbody>
<tr>
<td>Wheat</td>
<td>5/12</td>
<td>0</td>
<td>900–5184 (2640)</td>
</tr>
<tr>
<td>Rapeseed</td>
<td>5/17</td>
<td>0</td>
<td>2054–5646 (4027)</td>
</tr>
<tr>
<td>Maize</td>
<td>14</td>
<td>8</td>
<td>900–4251 (2997)</td>
</tr>
<tr>
<td>Pulses</td>
<td>10</td>
<td>3</td>
<td>900–6256 (3959)</td>
</tr>
<tr>
<td>Sugar beet</td>
<td>8</td>
<td>3</td>
<td>900–5598 (3630)</td>
</tr>
<tr>
<td>Rice</td>
<td>18</td>
<td>10</td>
<td>900–3761 (2817)</td>
</tr>
<tr>
<td>Soybean</td>
<td>13</td>
<td>10</td>
<td>900–3303 (2198)</td>
</tr>
<tr>
<td>Sunflower</td>
<td>15</td>
<td>6</td>
<td>900–3854 (2786)</td>
</tr>
<tr>
<td>Millet</td>
<td>12</td>
<td>10</td>
<td>900–4709 (2670)</td>
</tr>
<tr>
<td>Peanut</td>
<td>15</td>
<td>14</td>
<td>900–2711 (1788)</td>
</tr>
<tr>
<td>Cassava</td>
<td>22</td>
<td>15</td>
<td>900–2903 (1758)</td>
</tr>
</tbody>
</table>

* $T_{sow}$: temperature limit used for crop sowing; $T_b$: HU base temperature. Values are from Bondeau et al. (2007) or Waha et al. (2011) except for values in italic type, where the $T_b$ values are more common values found in the literature; PHU range used in this study with mean in brackets.
Table 2. LPJ-GUESS PFT harvest parameter values used in this study.

<table>
<thead>
<tr>
<th>PFT class</th>
<th>Harvest efficiency</th>
<th>Harvested parts</th>
<th>Harvest slow fraction</th>
<th>Slow fraction turnover rate (yr⁻¹)</th>
<th>Residue removal</th>
</tr>
</thead>
<tbody>
<tr>
<td>Trees</td>
<td>0.7</td>
<td>wood</td>
<td>0.33</td>
<td>0.04</td>
<td>1.0</td>
</tr>
<tr>
<td>Grass</td>
<td>0.5</td>
<td>leaves</td>
<td>0</td>
<td>1</td>
<td>1.0*</td>
</tr>
<tr>
<td>Crops</td>
<td>0.9</td>
<td>storage</td>
<td>0</td>
<td>1</td>
<td>0.75</td>
</tr>
</tbody>
</table>

* Natural grass 1.0, pasture and cover crop grass 0.0.
Table 3. Selected crop-intensive areas used in the FPAR-NDVI comparisons.

<table>
<thead>
<tr>
<th>Site</th>
<th>Position(^1) No. Cells</th>
<th>Cropland/pasture (%area)</th>
<th>Natural (\text{vegetation})^(^2) Crops(^3)</th>
<th>Irrigation</th>
<th>Seasonality</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Northern Africa</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Morocco</td>
<td>8.5° W, 32.5° N</td>
<td>6</td>
<td>41/27</td>
<td>C3g</td>
<td>Wh, Co, Pu</td>
</tr>
<tr>
<td>Algeria</td>
<td>6° E, 36° N</td>
<td>8</td>
<td>59/27</td>
<td>TeBE</td>
<td>Wh</td>
</tr>
<tr>
<td>Egypt</td>
<td>30.5° E, 30.5° N</td>
<td>6</td>
<td>58/26</td>
<td>C3g</td>
<td>Wh, Co, Ri</td>
</tr>
<tr>
<td><strong>Sahel</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Senegal</td>
<td>16.5° W, 13.5° N</td>
<td>8</td>
<td>57/16</td>
<td>C4g, TrBR</td>
<td>Mi, Pe</td>
</tr>
<tr>
<td>Burkina Faso</td>
<td>2.5° W, 12° N</td>
<td>10</td>
<td>28/27</td>
<td>TrBR, C4g</td>
<td>Mi</td>
</tr>
<tr>
<td>Sudan 1 + 2</td>
<td>33° E, 11.5° N/35° E, 13° N</td>
<td>10</td>
<td>66/17</td>
<td>TrBR, C4g</td>
<td>Mi, Pe, Wh(^4)</td>
</tr>
<tr>
<td><strong>East Africa</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Ethiopia</td>
<td>37° E, 5.5° N</td>
<td>15</td>
<td>35/23</td>
<td>TrBE, C3g</td>
<td>Wh(^3), Co, Mi, Pu</td>
</tr>
<tr>
<td>Uganda</td>
<td>31° E, 0.5° S</td>
<td>4</td>
<td>38/42</td>
<td>TrBE, TrBR</td>
<td>Pu, Mi, Co, Ma</td>
</tr>
<tr>
<td>Tanzania</td>
<td>38° E, 6° S</td>
<td>4</td>
<td>51/28</td>
<td>TrBE, TrBR</td>
<td>Co, Mi, Ma, Pu</td>
</tr>
<tr>
<td><strong>Southern Africa</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Botswana/South Africa</td>
<td>26.5° E, 24° S</td>
<td>8</td>
<td>40/30</td>
<td>C3g</td>
<td>Mi, Co, Wh</td>
</tr>
<tr>
<td>South Africa 1</td>
<td>28° E, 28.5° S</td>
<td>6</td>
<td>40/29</td>
<td>TrBE, C4g, MNE</td>
<td>Co, Wh, Sf</td>
</tr>
<tr>
<td>South Africa 2</td>
<td>18° E, 33.5° S</td>
<td>4</td>
<td>60/24</td>
<td>C3g</td>
<td>Co, Wh, Sf</td>
</tr>
</tbody>
</table>

\(^1\) The map reference refers to the south-west corner of the gridcell square; \(^2\) Dominant simulated natural vegetation plant functional types. TeBE: temperate broadleaved evergreen tree; TrBE: tropical broadleaved evergreen tree, TrBR: tropical broadleaved raingreen tree, MNE: mediterranean needle-leaved evergreen tree, C3g: C3 grass, C4g: C4 grass; \(^3\) Dominant crops (by reported area). Wh: wheat (+barley etc.), Mi: millet (+Sorghum), Co: corn(maize), Ri: rice, Pe: peanut(groundnuts), Ma: maniok(cassava), Pu: pulses, Sf: sunflower; \(^4\) wheat not simulated because of temperature restriction.
Table 4. Setup of the different LPJ-GUESS simulations.

<table>
<thead>
<tr>
<th></th>
<th>Description</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>Natural vegetation only, variable climate and CO₂.</td>
</tr>
<tr>
<td>2</td>
<td>Standard run: variable climate, CO₂ and land cover.</td>
</tr>
<tr>
<td></td>
<td>Cover crop grass. Standard harvest parameters.</td>
</tr>
<tr>
<td>3</td>
<td>As (2), but with spinup land cover (1901 data).</td>
</tr>
<tr>
<td>4</td>
<td>As (2), but with spinup climate (1901–1930 data cycled).</td>
</tr>
<tr>
<td>5</td>
<td>As (2), but with spinup temperature (1901–1930 data cycled).</td>
</tr>
<tr>
<td>6</td>
<td>As (2), but with spinup precipitation (1901–1930 data cycled).</td>
</tr>
<tr>
<td>7</td>
<td>As (2), but with spinup radiation (1901–1930 data cycled).</td>
</tr>
<tr>
<td>8</td>
<td>As (2), but with spinup CO₂ (1901 value).</td>
</tr>
<tr>
<td>9</td>
<td>As (2), but without cover crop grass.</td>
</tr>
<tr>
<td>10</td>
<td>As (2), but with no crop residue removal.</td>
</tr>
<tr>
<td>11</td>
<td>As (2), but with 100% crop residue removal.</td>
</tr>
</tbody>
</table>
Fig. 1. Crop phenology in LPJ-GUESS. *HU sum: heat unit sum (dynamic potential HU adapted to local climate); LAI: leaf area index; HI: harvest index.
Fig. 2. Location of selected cropland areas. Cells with cropland fractions > 25% are indicated in green. The climate seasonality types for 1990 used for sowing date calculation by LPJ-GUESS are indicated by hatched grey. The two Sudan crop regions are treated as a single region in the text. The seasonality types are defined in the Methods section.
Fig. 3. Example of land use change modelling at a moist tropical site (9°W, 5.5°N) during the period 1901–2006 using the Ramankutty and Foley (1999) land cover dataset. (A) Carbon fluxes to the atmosphere associated with crop harvest and land use change (cropland expansion/deforestation). The “slow C pool” represents the oxidation of carbon from the long-lived wood pool harvested at cropland expansion. (B) Lifetime of natural vegetation stands. New stands are created at cropland abandonment and destroyed at subsequent cropland expansion. (C) Cropland area fraction, showing the expansion and reduction of cropland area. (D) NPP of three PFTs at one of the natural stands created at cropland abandonment, representing a typical plant successional sequence where an initial dominance of grass is followed by forest, dominated first by deciduous and later by evergreen trees.
Fig. 4. Standardised anomalies of monthly observed NDVI and FPAR modelled by LPJ-GUESS with and without landuse functionality, relative to the gridcell mean and standard deviation, averaged over latitude bands for the period 1982–2006.
Fig. 5. Standardised seasonal anomalies (average for 4–15 adjacent cells, see Table 3 for details) of NDVI and FPAR, modelled by LPJ-GUESS with and without landuse (LU) functionality, relative to the African total mean and standard deviation, for the period 1982–2006. “LPJ-GUESS FPAR” represents potential natural vegetation only, while “LPJ-GUESS LU FPAR” represents potential natural vegetation plus cropland and pasture. For the Moroccan and Egyptian sites, an alternative sowing regime for irrigated crops is indicated (precipitation-dependent sowing).
Fig. 6. Difference in distance index (di) for monthly observed NDVI and modelled FPAR by adding land use functionality with cropland and pasture. Cropland is simulated without cover crop grass (negative values are improvements to the FPAR vs. NDVI fit).
Fig. 7. Standardised mean anomalies for the Sahel region of yearly maximum observed NDVI and FPAR, modelled by LPJ-GUESS with and without landuse functionality, and precipitation for the period 1982–2006.
Fig. 8. Scatterplots of LPJ-GUESS modelled yield versus FAO reported yield of the seven most widespread crops in Africa (1991–1995 country means). The five countries with the largest area of each crop are indicated in red.
Fig. 9. (A) Time series of LPJ-GUESS modelled irrigated, rainfed and average yields and reported FAO yields for maize 1971–2005 for selected African countries. (B) Scatter plot of LPJ-GUESS modelled versus FAO reported maize yields 1971–2005 with linear regression lines.
Fig. 10. Modelled cumulative Net Ecosystem Carbon Balance (NECB) for the African continent, with alternative landuse settings: natural vegetation only (“GUESS”), natural plus cropland and pasture vegetation with standard settings (75% crop residue removal, cover crop grass growth, “GUESS-crop”), without crop residue removal (“GUESS-crop no res rem”) and without cover crop grass (“GUESS-crop no cc grass”). Positive values represent a flux to the atmosphere.
Fig. 11. (A) Separate contributions of CO$_2$, climate and land use change to the modelled 1901–2006 cumulative Net Ecosystem Carbon Balance (NECB) for the African continent, derived by comparing a simulation (“Net”) forced by historical CO$_2$, climate and land use with separate simulations where one of these components was kept at the spinup levels throughout the simulation period. For example, the climate component (“Climate”) was derived by subtracting the cumulative NECB of a simulation that considered only CO$_2$ and land use change from the “Net” simulation that considered CO$_2$, land use and climate change. (B) Separate contributions of changing temperature, precipitation and radiation to the modelled 1901–2006 cumulative African NECB, derived by comparing the “Net” simulation above with separate simulations where one climate component was kept at the spinup levels throughout the simulation period as in (A). “Climate” is the net climate component as in (A). Positive values represent a flux to the atmosphere.
Fig. 12. (A) Difference in modelled Net Ecosystem Carbon Balance (NECB) between LP-GUESS with and without land use functionality (cropland and pasture). (B) Carbon fluxes associated with landuse change (not including carbon added to litter). The values are the means for the period 1961–1990. Positive values represent a flux to the atmosphere.