

the change in the net radiative flux anomaly versus the change in the surface air or the sea surface temperature anomaly (Trenberth et al., 2010; Dessler, 2010). However, Spencer and Braswell (2010) suggested that such a method would be ambiguous considering its failure to separate the changes in the net radiative flux anomaly due to feedback (caused by temperature changes) from changes due to varying radiative forcing through, for example, changing clouds. Masters (2012) found that the method used by Dessler (2010) generated large uncertainties causing difficulties in drawing conclusions from the results.

The climate feedback parameter is closely related to the equilibrium climate sensitivity (Gregory et al., 2004). For example, a true value of this parameter of $1.2 \text{ W m}^{-2} \text{ K}^{-1}$, which is a typical value from climate model simulations (Gregory et al., 2004), corresponds to a climate sensitivity of $3.7/1.2 = 3.1 \text{ K}$ for a doubling of the carbon dioxide mixing ratio.

An unusually high value of the climate feedback parameter of $6 \text{ W m}^{-2} \text{ K}^{-1}$ is suggested by the phase plane plots in Spencer and Braswell (2010). This corresponds to a very low climate sensitivity that disagrees with the majority of the other estimations of the climate sensitivity (Knutti and Hegerl, 2008; Randall et al., 2007; Huber et al., 2011). A discussion of the various methods for estimation of the climate sensitivity is beyond the scope of this work. Here we discuss a method for estimating the value of the climate feedback parameter from satellite radiative flux data and leave the question how to relate the result from this method to the equilibrium climate sensitivity to future work.

In Trenberth et al. (2010, 2011) and Dessler (2011) the authors draw special attention to the role of El Niño Southern Oscillation (ENSO) during the first decade of this millennium. Loeb et al. (2012) have discussed how TOA radiative flux anomalies have changed due to ENSO events.

In this work it is suggested from the analysis of the relation between the surface air temperature anomaly and the TOA net radiative flux anomaly, derived from the Clouds and Earth's Radiant Energy Systems (CERES) Energy Balanced and Filled (EBAF)

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product (Loeb et al., 2009), that an assumption that ENSO has completely dominated the temperature changes during the whole period September 2000–May 2011 is not valid. However, it is suggested as a hypothesis that ENSO dominated the change in the surface air temperature anomaly in parts of this period especially the five year period mid-2006 to mid-2011. The study of that period appears to be useful for estimating a value of the climate feedback parameter by using phase plane plots similar to those suggested by Spencer and Braswell (2010).

2 Theoretical background

According to the concepts of radiative forcing and feedback, as described by Gregory et al. (2004), the net downward radiative flux at the top of the atmosphere (TOA) is separated into a forcing term, positive downwards, and a feedback term, positive upwards.

$$N = F - H \quad (1)$$

Gregory et al. (2004) adopted results from model experiments with Global Circulation Models (GCMs) that the feedback term is proportional to the surface air temperature change:

$$H = \alpha \Delta T \quad (2)$$

This gives the following simple linear model equation for the TOA radiative flux:

$$N = F - \alpha \Delta T \quad (3)$$

Gregory et al. (2004) used this simple model to study feedback in Global Circulation Models (GCMs) and coupled Atmospheric Ocean GCMs (AOGCMs) in model experiments with $2 \times \text{CO}_2$ och $4 \times \text{CO}_2$. In such experiments the radiative forcing in principle is changed instantaneously and after that kept constant.

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Gregory et al. (2004) plotted the changes in the TOA net downward radiative flux N vs. the surface air temperature changes. The relaxation process in the model experiments resulted in straight lines, the slopes of which correspond to the feedback parameter α , the intercepts on the y-axis give values of the radiative forcing and the intercepts on the x-axis are measures of the climate sensitivity in case of a $2\times\text{CO}_2$ model experiment.

Such plots were done for total net radiative flux, for long wave (LW) and for short wave (SW) flux.

Spencer and Braswell (2010) used the same conceptual framework and model equation for the net radiative flux for evaluating radiative flux observations with satellites. They reasoned that if satellite data of net radiative flux anomalies are plotted vs. temperature anomalies and the points are connected in time sequence to obtain a phase plane plot, segments of straight lines should appear in periods when the radiative forcing has a constant value. The slope of such segments would be equal to the feedback parameter α . If segments with the same slope would appear several times in the phase plane plot at difference to more randomly occurring slopes of other line segments this would be a strong indication that this slope corresponds to the feedback parameter value.

Spencer and Braswell (2010) used one month averaging and three month averaging for CERES data. For data from The Earth Radiation Budget Experiment (ERBE) they used 216 days averages. They found indications of the expected segments of straight lines (also called striations) when using phase plane plots of the net radiative flux anomalies vs. middle troposphere temperature anomalies both using ERBE and CERES data. They selected the middle troposphere temperature because it is more correlated to the TOA radiative flux than the surface air temperature.

For the application of the method by Spencer and Braswell (2010) it is essential to use anomalies of the net radiative flux and the temperature in order to eliminate the seasonal variations. This is because the time periods are too short in order to plot only whole years as in the case of Gregory et al. (2004).

Spencer and Braswell (2010) also found interesting great loops in the phase plane diagrams, with large excursions of the temperature anomaly, especially in connection with the Mount Pinatubo cooling in ERBE data and with the 2008 La Niña cooling and subsequent warming in CERES data.

They further applied their phase plane method to the radiative flux data from AOGCMs where they studied phase planes not only with total net radiative flux but also with LW and SW fluxes. In those cases they used surface air temperature anomalies and 11 months low pass filtered averages. They studied periods of 50 yr of length.

Indications of striations were found in the LW phase planes of four of the AOGCMs and their slopes coincided with the feedback parameter values determined by Forster and Taylor (2006). However, in the phase plane diagrams for total net radiative flux or SW flux none of the AOGCMs produced significant striations. The result with LW phase planes is a support for the method since the AOGCMs confirm that if many striations appear with a common slope, this slope equals the value of the feedback parameter.

The present study is based on the idea of using phase plane plots according to Spencer and Braswell (2010). The main innovation in this study is the usage of phase plane plots where the radiative flux with a time lag is plotted vs surface temperature.

3 Data and methods

The TOA radiative flux data used was retrieved from the CERES EBAF product available online (NASA CERES EBAF, 2012). More details about the radiative flux data and the calculation of anomalies are given in the supplementary material. Loeb et al. (2009) give a background of and discuss the design of the CERES EBAF product. The surface air temperature anomalies used are from HadCRUT3 (Brohan et al., 2006).

The monthly anomalies for the net radiative fluxes from CERES EBAF were calculated as deviations from the monthly averages for the period 2001–2011. Both the net radiative flux anomalies and the surface air temperature anomalies were smoothed by calculating running centered 13 months averages. The choice of 13 months for

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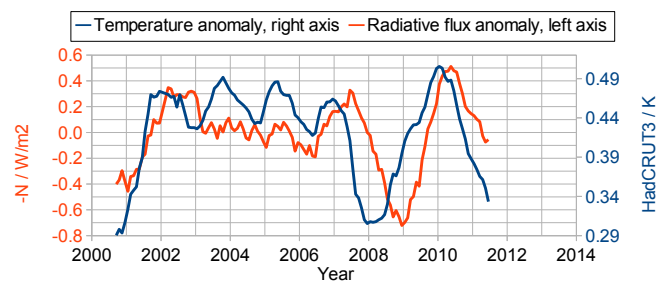


Fig. 1. Thirteen months running averages of net radiative flux anomalies and surface air temperature anomalies plotted as functions of time.

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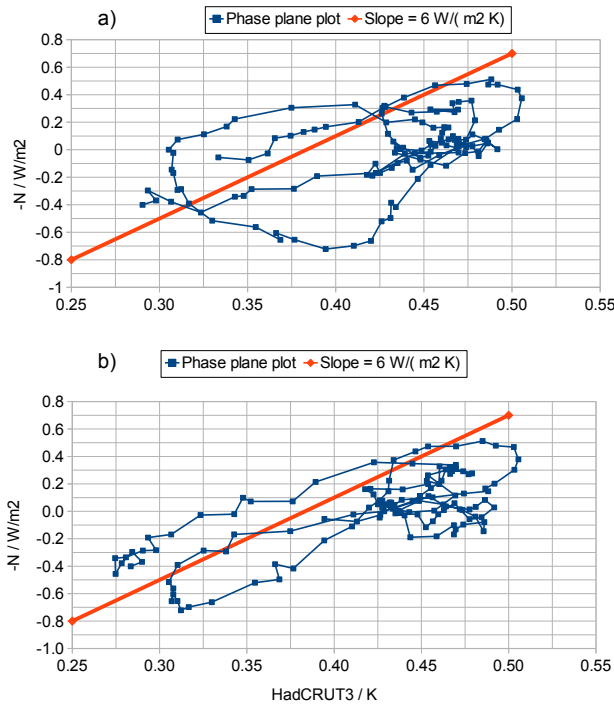


Fig. 2. Phase plane plots of the same data as in Fig. 1. Net radiative flux anomalies are plotted vs surface air temperature anomalies. In (a) there is no lag while in (b) the net radiative flux anomalies lag the temperature anomalies with seven months.

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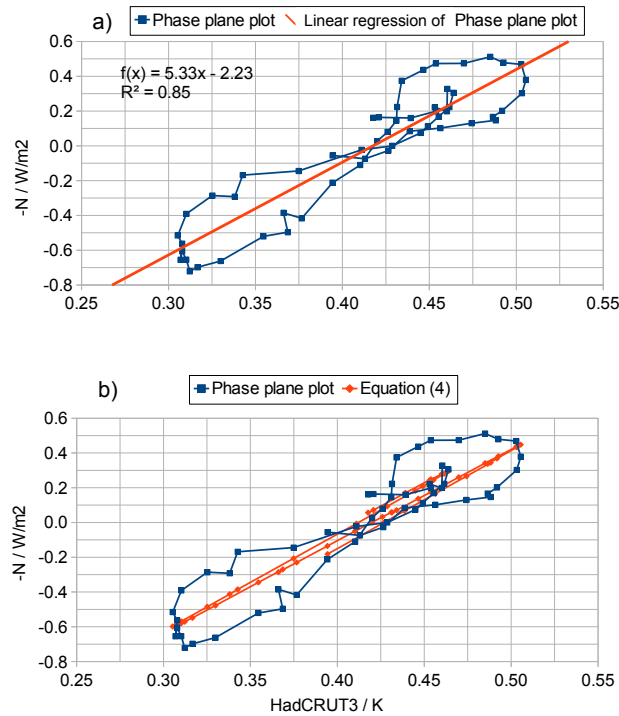


Fig. 3. Phase plane plots like in Fig. 2b with seven months lag but for the time interval mid-2006 to mid-2011. In (a) with constant forcing F according to Eq. (5). In (b) a possible linear trend in the forcing has been considered according to Eq. (4).

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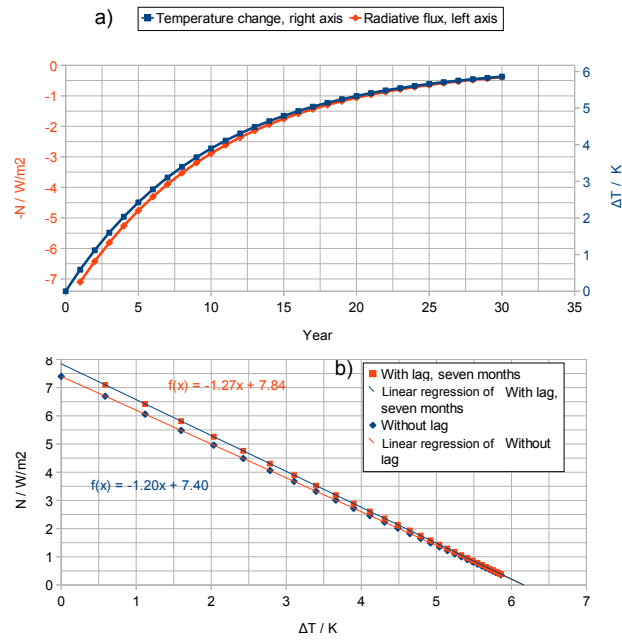


Fig. 4. (a) Net radiative flux change and temperature change plotted as a function of time for a $4\times\text{CO}_2$ experiment with a lag in the net radiative flux change as explained in the text. **(b)** Phase plane plots of a $4\times\text{CO}_2$ experiment with and without a lag in the net radiative flux change.